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Author: Brownstein, Chase D.

Source: Journal of Paleontology, 92(6) : 1115-1129

Published By: The Paleontological Society

URL: <https://doi.org/10.1017/jpa.2018.42>

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The distinctive theropod assemblage of the Ellisdale site of New Jersey and its implications for North American dinosaur ecology and evolution during the Cretaceous

Chase D. Brownstein

Stamford Museum and Nature Center, Stamford CT <chasethedinosaur@gmail.com>

Abstract.—The Cretaceous landmass of Appalachia has preserved an understudied but nevertheless important record of dinosaurs that has recently come under some attention. In the past few years, the vertebrate faunas of several Appalachian sites have been described. One such locality, the Ellisdale site of the Cretaceous Marshalltown Formation of New Jersey, has produced hundreds of remains assignable to dinosaurs, including those of hadrosauroids of several size classes, indeterminate ornithopods, indeterminate theropods, the teeth, cranial, and appendicular elements of dromaeosaurids, ornithomimosaurs, and tyrannosauroids, and an extensive microvertebrate assemblage. The theropod dinosaur record of the Ellisdale site is currently the most extensive and diverse known from the Campanian of Appalachia. Study of the Ellisdale theropod specimens suggests that at least four or more non-avian theropod taxa are represented at the site, including tyrannosauroids, ornithomimosaurs, several different morphotypes of dromaeosaurids that are the first of that clade described from New Jersey, and indeterminate theropods. The specimens are important for increasing current knowledge about the theropod diversity of the Atlantic Coastal Plain (ACP) during the Campanian by representing the most speciose assemblage of the group during the time in the ACP as well as for shedding light on Appalachian dinosaur ecology and biogeography generally.

Introduction

The landmass of Appalachia was created when the Western Interior Seaway flooded the interior of North America and separated the eastern portion of the continent from the west midway through the Cretaceous (Roberts and Kirschbaum, 1995; Russell, 1995; Schwimmer, 1997, 2002; Sampson et al., 2010). The fossil record of dinosaurs from the portion of eastern North America that corresponds to Appalachia is poor (e.g., Weishampel and Young, 1996; Schwimmer, 1997; Weishampel, 2006), but the publication of several new taxa (e.g., Carr et al., 2005; Prieto-Márquez et al., 2016a, 2016b) and the first reports of some clades of dinosaur (e.g., ceratopsians) from this landmass (Longrich, 2016; Farke and Phillips, 2017) in the past several years have greatly increased our knowledge of the obscure fauna of the landmass.

The discovery of several new sites in the eastern United States has also illuminated the vertebrate faunas of Appalachia (e.g., Weishampel, 2006; Denton et al., 2011; Schwimmer et al., 2015). Study of these faunas has had broad implications for modeling the evolution of several clades of dinosaurs, generally suggesting that Appalachia harbored relict forms isolated from more derived relatives on other landmasses and continents by the Western Interior Seaway (e.g., Weishampel and Young, 1996; Schwimmer, 1997; Carr et al., 2005; Brusatte et al., 2011; Denton et al., 2011; Schwimmer et al., 2015; Prieto-Márquez et al., 2016a, 2016b). Examples include the Stokes Quarry site of South Carolina, from which teeth and

other elements attributed to the tyrannosauroid *Appalachiosaurus montgomeriensis* Carr, Williamson, and Schwimmer, 2005, the dromaeosaurid *Saurornitholestes langstoni* (Sues, 1978), and one or two other dromaeosaurid morphotypes have been described (Schwimmer et al., 2015). Other important sites include Phoebe Landing in North Carolina, which has preserved teeth and postcranial elements comparable to *Dryptosaurus aquilunguis* (Cope, 1866) and other tyrannosauroids, *Lophorhynchon*, and *Hadrosaurus*, as well as the bones of ornithomimosaurs and the holotype of the gigantic hadrosauroid *Hypsiobema crassicauda* (Cope, 1869) (e.g., Miller, 1967; Baird and Horner, 1979; Weishampel and Young, 1996; Schwimmer, 2016), and the Hannahatchee site of Georgia, which has preserved teeth and postcranial remains assigned to *Appalachiosaurus montgomeriensis*, indeterminate ornithomimosaurs, and indeterminate hadrosauroids (e.g., Schwimmer et al., 1993). These sites have also been noted for their microvertebrate remains, including those of mammals, small squamates, and amphibians, making such localities extremely significant to research regarding faunal changes in North America during the Late Cretaceous (e.g., Grandstaff et al., 1992; Gallagher, 1993, 1997; Weishampel and Young, 1996; Denton et al., 2004, 2011).

One of the most important Late Cretaceous fossil sites from the eastern United States, an outcrop of the middle to late Campanian Marshalltown Formation (which is equivalent in age to the upper Tar Heel and Donoho Creek formations of the Carolinas; e.g., Miller et al., 2004; Harris and Self-Trail, 2006),

is the Ellisdale site of New Jersey, discovered in 1980 by Robert C. O'Neill and Robert K. Denton Jr. (e.g., Weishampel and Young, 1996). Since then, tens of thousands of fossil specimens have been recovered from Ellisdale pertaining to a menagerie of vertebrate and invertebrate clades (e.g., Weishampel and Young, 1996; Denton et al., 2004, 2011). The site has yielded an extremely important record of microvertebrates, with finds including some of the best specimens of mammals from the Late Cretaceous of eastern North America (Grandstaff et al., 1992; Gallagher, 1993; Denton et al., 2011), a species of batrachosauroidid salamander (*Parrisia neocaesariensis* Denton and O'Neill, 1998), and a species of teiid lizard (*Prototeius stageri* Denton and O'Neill, 1995). Indeed, the site has been noted for its importance in understanding the biodiversity of coastal Appalachian ecosystems (e.g., Grandstaff et al., 1992; Denton and O'Neill, 1995, 1998; Denton et al., 2004, 2011; Weishampel, 2006). Dinosaur remains from Ellisdale, particularly those of ornithopods, have been recovered in the hundreds (Denton et al., 2011; personal observation, 2017). These fossils preserve both skull and appendicular elements from several species, with the taxa *Dryptosaurus*, *Hadrosaurus*, and *Hypsibema* reported from the site along with the remains of indeterminate theropods and hadrosaurids (e.g., Grandstaff et al., 1992; Gallagher, 1993). However, the sheer number of dinosaur specimens collected from Ellisdale warrants further study of specimens identifiable past such general ranks as "Theropoda indet." in order to better characterize dinosaur diversity at this site and on Appalachia generally.

Here, theropod specimens from Ellisdale referable to more inclusive clades within that group are described. Several recent studies have noted that theropod teeth may not be diagnostic to the species level (e.g., Longrich, 2008a; Sankey, 2008; Williamson and Brusatte, 2014). Nevertheless, some studies have quantified differences in apparently distinct theropod tooth morphotypes (e.g., Larson and Currie, 2013; Williamson and Brusatte, 2014). Caution was taken with assigning any of the specimens to particular genera. The Ellisdale site theropod specimens, which represent a diverse assemblage of theropod dinosaur clades, are important for further illuminating the diversity of theropod dinosaurs in the Atlantic Coastal Plain and on Appalachia generally, for their biogeographic implications for dinosaur provincialism on Appalachia, and for understanding the differences between Appalachian and Laramidian faunas.

Geological setting

The Englishtown Formation, made up of slightly glauconitic lignitic cross-bedded sands (e.g., Owens and Sohl, 1969; Owens et al., 1998), crops out at the base of the exposed deposit at Ellisdale and is uncomfortably overlain by a series of highly lignitic bedded clays, flaser sands, and thin layers of cross-bedded sands (e.g., Grandstaff et al., 1992). These latter three sand layers are overlain by marine sands and glauconitic sediments, and all have been identified as the Marshalltown Formation (e.g., Grandstaff et al., 1992; Denton and O'Neill, 1995; Denton et al., 1998, 2004, 2011). The Marshalltown Formation was dated at ca. 75.7–71.2 Ma by Miller et al. (2004). More recently, palynological study of the bedded clays that straddle

the main fossiliferous layer of siderite clay pebble-filled sand (Grandstaff et al., 1992) has revealed their age as between 76.4–79.6 Ma (e.g., Denton and Tashjian, 2012), placing the Marshalltown Formation in the middle to late Campanian and thus making the unit closely comparable temporally with well-documented western deposits such as the Kaiparowits, Dinosaur Park, Judith River, and Two Medicine formations (e.g., Gates et al., 2010), as well as the upper Tar Heel and Coachman formations of the Carolinas (e.g., Harris and Self-Trail, 2006; Schwimmer et al., 2015). At least four different paleoenvironments are represented at Ellisdale based on the fauna and geology of the site: lagoonal/backbay, estuarine, marine, and terrestrial (e.g., Grandstaff et al., 1992; Denton and O'Neill, 1995; Denton et al., 2004, 2011; Denton and Tashjian, 2012), and several models have been proposed to account for the apparent pre-burial transport and disarticulation of species from these several environments at the Ellisdale site. For example, the site has been interpreted as a storm deposit (e.g., Gallagher et al., 1986; Tajishan, 1990; Grandstaff et al., 1992). Denton et al. (2004) proposed that the fossiliferous deposits at Ellisdale were formed when coastal storms flooded a barrier island and washed, churned, and disarticulated the bodies of dead marine and island-inhabiting organisms trapped in the storm surge and washed them into an inlet, where they mixed with the eroded, river-deposited remains of species from environments further upland. More recently, Denton and Tashjian (2012) suggested that formation of the layers at the Ellisdale site occurred in several steps in a backbay/estuarine setting. The first step, deposition and regression of sediments from the Englishtown Formation, was followed by transgressions that rapidly changed environments surrounding the backbay and thus changed the type of flora and fauna being deposited (Denton and Tashjian, 2012). The deposition of species from different terrestrial and coastal environments, in this model, would have been interspersed by storm events that mixed together already deposited fossil beds and dead biological matter from greater distances (Denton and Tashjian, 2012). The hypotheses of Denton et al. (2004) and Denton and Tashjian (2012) are supported by the presence of heavy erosion on some of the theropod remains and only very slight signs of abrasion on others. The Marshalltown Formation at Ellisdale is overlain by the early Maastrichtian Mount Laurel Formation and Pleistocene gravel deposits (e.g., Grandstaff et al., 1992; Miller et al., 2004).

Materials and methods

Referral of teeth to specific dinosaur clades.—In order to support referral of the teeth described to particular clades, they were carefully inspected for phylogenetically informative characters included in the matrices of Carr et al. (2017) (for the tyrannosauroid teeth) and Turner et al. (2012) (for the dromaeosaurid teeth). Principle component and discriminant function analyses (PCA and DFA) were not undertaken on the Ellisdale theropod teeth due to the fragmentary or eroded nature of many of the specimens. Nevertheless, the teeth are informative enough to be assigned to specific clades, and possibly genera, of theropod dinosaurs.

Measurement and anatomical methodology.—The proposed standardized nomenclature for theropod teeth described by Hendrickx et al. (2015) was followed. Tooth dimensions are described using: (1) crown height (CH), measured from the middle of the base of the tooth to the apex; (2) crown base length (CBL), measured from the middle of the mesial to the middle of the distal faces of the tooth; and (3) crown base width (CBW), measured at the greatest labiolingual width of the base of the tooth. Because many of the Ellisdale specimens are fractured and all are very fragile, measurements of small teeth and their denticles were taken using digital calipers and rounded to the nearest tenth of a millimeter.

Institutional abbreviation.—NJSM, collections of the New Jersey State Museum, New Jersey, USA.

Systematic paleontology

Theropoda Marsh, 1881
 Tetanurae Gauthier, 1986
 Coelurosauria von Huene, 1914
 Tyrannosauroida Osborn, 1905
 Tyrannosauroida indet.
 Figures 1, 2.1–2.5

Description.—NJSM 14682 is a slightly eroded pedal phalanx of a tyrannosauroid dinosaur measuring 113 mm proximodistally and has a circumference of 123 mm at its diaphysis. This bone is identified as a phalanx III-1 or III-2 because it is not proximodistally elongate and mediolaterally compressed (conditions in the pedal phalanges of digit II) or proximodistally compressed (as in the pedal phalanges of digit IV or III-3). The

phalanx is robustly built, with a strongly arched ventral surface in medial and lateral view. The collateral ligament pits are both deep and teardrop-shaped, both reaching their deepest point proximally. The lateral collateral ligament pit is better preserved than the medial. The proximal and distal surfaces are moderately worn, yet the preserved bone indicates that only a very slight sulcus separated the distal hemicondyles. On the dorsal surface, a deepened, teardrop-shaped fossa sits just proximal to the distal hemicondyle, identified as the extensor fossa for the insertion of the *M. extensor digitorum longus* (Carrano and Hutchinson, 2002). The shaft of the bone is arched dorsoventrally and pinched mediolaterally between the mediolaterally expanded proximal and distal ends of the bone. The middle portion of the bone is slightly less abraded than the proximal and distal ends, the former semi-rectangular in proximal view due in part to erosion. In medial and lateral views, the dorsal surface of the shaft is flattened. In lateral and medial views, the shaft expands dorsoventrally towards the proximal end, unlike the pinched morphology seen in dorsal view. The shaft is slightly rounded dorsally, ventrally, medially, and laterally. Towards the ventral part of the proximal most portion of the lateral and medial surfaces of the bone, there are slight depressions. The proximal surface, like the distal, is eroded. However, enough bone is preserved to show the presence of a lip around the proximal articular surface with its most distally outstretched point at the dorsal apex of the bone. The proximal surface is semi-rectangular in shape, and there is no vertical inflection on the proximal articular facet. Measurements of this bone may be found in Table 1.

NJSM 14682 is most parsimoniously placed within Tyrannosauroida based on several features of the bone. First, the bone's dimensions are much larger than those of definite ornithomimosaur phalanges known from the Late Cretaceous of the Atlantic Coastal Plain (e.g., AMNH FARB 2553, NJSM 14686) (Table 2). NJSM 14682, at 113 mm proximodistally, is also longer than the largest of any of the pedal phalanges of *Gallimimus* by more than 10 mm (e.g., Osmólska et al., 1972, table 2). Thus, the bone is appreciably larger than the corresponding element in any Cretaceous ornithomimid (e.g., Makovicky et al., 2004; Kobayashi and Barsbold, 2005a, table 6.3; Longrich, 2008b, fig. 13) excluding *Deinocheirus* and *Beishanlong*, which have pedal phalanges of similar length (e.g., Makovicky et al., 2009, fig. 3; Lee et al., 2014, extended data fig. 5). However, the author regards it as unlikely that NJSM 14682 comes from a deinocheirid, given that the group is currently only represented by Cretaceous genera from Asia (e.g., Lee et al., 2014). NJSM 14682 may further be distinguished from those of ornithomimids by the lack of proximally projecting ridges that extend past the proximal end of the bone and form a concavity on the ventral surface, a robust distal end that has approximately the same mediolateral width as the proximal (Table 1), having a deeper flexor fossa relative to the pedal phalanx III-1 of ornithomimosaurs, having a deeply concave proximal surface, and having a mediolaterally “pinched,” as opposed to gently arched, diaphysis (e.g., Osborn, 1921, fig. 3a; Osmólska et al., 1972, fig. 17; Kobayashi and Barsbold, 2005a, fig. 6.10G, 6.10H, 2005b, fig. 17B; Choiniere et al., 2012, fig. 14; Cullen et al., 2013, figs. 2, 3; McFeeters et al., 2016, fig. 11; Tsogtbaatar et al., 2017, fig. 4). NJSM

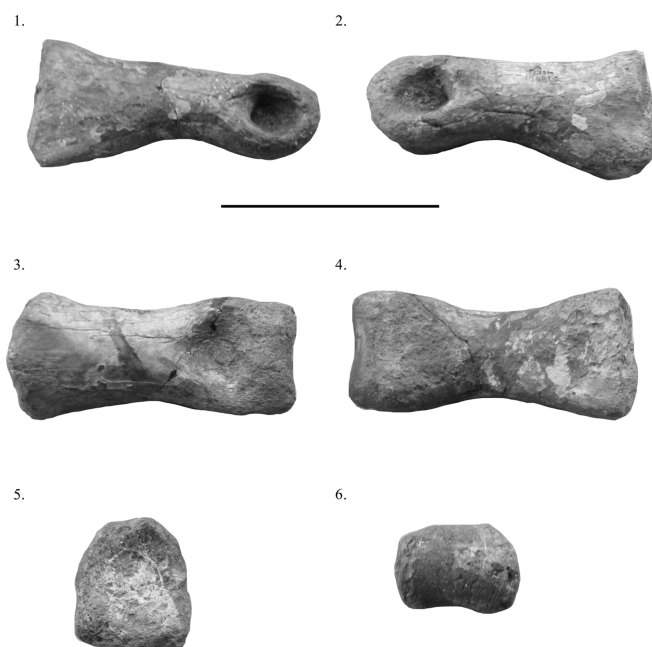


Figure 1. Tyrannosauroid pedal phalanx NJSM 14682 in lateral (1), medial (2), dorsal (3), ventral (4), proximal (5), and distal (6) views. Scale bar = 50 mm.

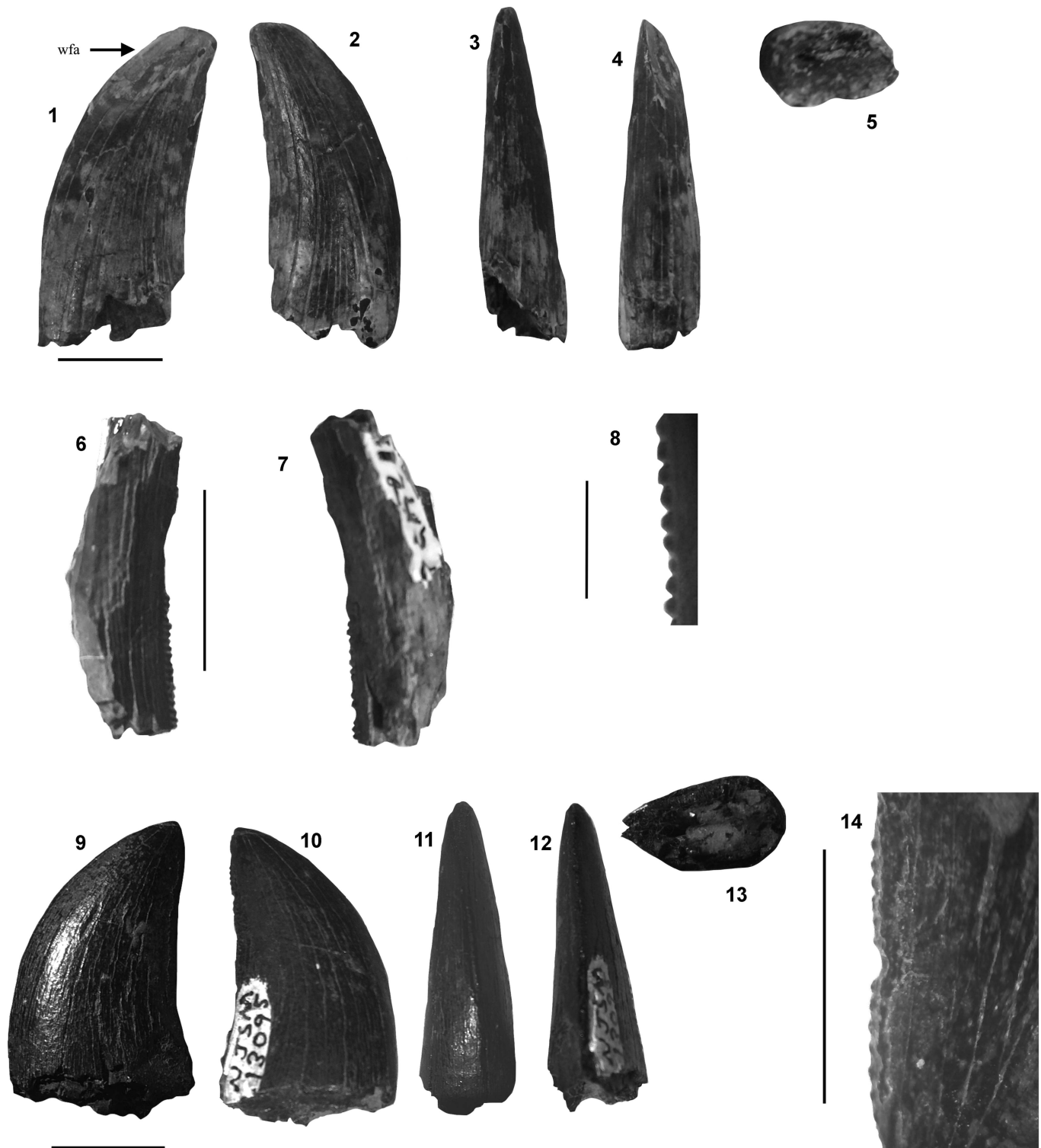


Figure 2. cf. *Dryptosaurus* maxillary or dentary teeth NJSM 16601 (1–5), NJSM 12436 (partial tooth) (6–8), and NJSM 13095 (9–12) in labial (1, 6, 9), lingual (2, 7, 10), mesial (3, 11), distal (4, 12), and basal (5, 14) views, with magnified views of NJSM 12436 and NJSM 13095 (H, N); wfa = wear facet. Scale bar = 10 mm (1–7, 9–13), = 5 mm (8, 14).

14682 is also more mediolaterally widened than the pedal phalanges of dromaeosaurids and troodontids and comes from an animal of much greater size than any known from the former group known from eastern North America or from the latter group entirely (e.g., Kiernan and Schwimmer, 2004; Makovicky

and Norell, 2004; Norell and Makovicky, 2004; Schwimmer et al., 2015). NJSM 14682 may additionally be distinguished from dromaeosaurids, troodontids, and therizinosaurs on the basis of the lack of a deepened sulcus separating the distal hemicondyles (e.g., Clark et al., 2004; Makovicky and Norell,

Table 1. Measurements for Ellisdale theropod specimens (in mm).

Element	Assignment	Proximodistal length	Mediolateral width (proximal)	Mediolateral width (distal)	Dorsoventral width (proximal)	Dorsoventral width (distal)	Circumference (diaphysis)
NJSM 14682	Tyrannosauroida indet.	113	47	43	53	33	125
NJSM 14686	Ornithomimosauria indet.	67	31	25	25	19	74

Table 2. Measurements for Ellisdale theropod teeth (in mm).

Specimen	Assignment	CH	CBL	CBW	distal denticles/5 mm	mesial denticles/5 mm	CBW/CBL
NJSM 16601	Tyrannosauroida indet.	35	13	7	n/a	n/a	0.54
NJSM 12436 (partial crown)	cf. <i>Dryptosaurus</i> (morphotype A)	27.5	n/a	6	12	n/a	n/a
NJSM 13095	cf. <i>Dryptosaurus</i> (morphotype A)	27	15	8	?12	n/a	0.53
NJSM 13734	Tyrannosauroida indet. (morphotype B)	25	n/a	6	17	n/a	n/a
NJSM 14158 (larger)	Dromaeosauridae indet. (morphotype A)	25	10	5	16	19	0.5
NJSM 14158 (smaller)	Dromaeosauridae indet. (morphotype A)	25	14	6+	n/a	n/a	0.43 +
NJSM 12436 (complete crown)	Dromaeosauridae indet. (morphotype A)	26	16+	8	n/a	n/a	0.5
NJSM 14404 (larger)	Dromaeosauridae indet. (morphotype B)	n/a	n/a	n/a	n/a	n/a	n/a
NJSM 14404 (smaller)	Dromaeosauridae indet. (morphotype B)	n/a	n/a	n/a	~ 17 + (6 d/1.8 mm)	n/a	n/a
NJSM 16623	Theropoda indet.	16	7	2	n/a	n/a	0.29
NJSM 15319	Theropoda indet.	5	0.5	0.25	n/a	n/a	0.50
NJSM 16611	Dromaeosauridae indet. (morphotype A)	16	8.0	4.0	n/a	n/a	0.50

2004; Norell and Makovicky, 2004). NJSM 14682 is also considerably more elongate than the extremely proximodistally shortened pedal phalanges of some therizinosaurs (e.g., Clark et al., 2004). However, large size, a mediolaterally “pinched diaphysis” on III-2, robust morphology, a deepened proximal articular facet, a subtle sulcus separating the distal hemicondyles, and deepened collateral ligament pits are also found in tyrannosauroids, especially the larger species of the Late Cretaceous (e.g., Table 2; Lambe, 1917, fig. 49; Carr and Williamson, 2000, figs. 3, 4, 8, 15; Holtz, 2004; Farlow et al., 2013). Indeed, the deep hyperextensor pit on NJSM 14682 unites the bone with the corresponding element in *Appalachiosaurus montgomeriensis* (Red Mountain Museum specimen 6770; personal communication, R.K. Denton, 2017). Thus, it may be confidently referred to Tyrannosauroida based on the aforementioned presence of features.

Several morphologies on the bone, however, suggest against it being a tyrannosaurid. The ratio of the mediolateral width of the diaphysis to the proximodistal length of NJSM 14682 is greater than 3.0 (= 3.3) (Table 2) (e.g., Brusatte et al., 2010; Brusatte and Carr, 2016), the hyperextensor pit is deepened, unlike tyrannosaurids, where it is shallow, but as in *Appalachiosaurus montgomeriensis* (e.g., Lambe, 1917, fig. 49; Carr and Williamson, 2000, figs. 3, 4, 8, 15; Holtz, 2004; Farlow et al., 2013; personal communication, R.K. Denton, 2017). The phalanx could only be coded for one character in the Theropod Working Group Matrix (Brusatte et al., 2014), and thus was not included in a phylogenetic analysis. However, the morphology of the specimen makes it confidently assignable to Tyrannosauroida.

NJSM 16601 (Fig. 1.7–1.11) is the eroded tooth of a large theropod dinosaur. The tooth measures 13 mm long mesiodistally and 4 mm wide labiolingually at its base. The tooth’s CH is 33 mm, and the specimen does not preserve denticles. The tooth is worn, with several vertical cracks in the enamel visible in labial and lingual views. The specimen bears a striation-filled

elliptical wear facet, 10 mm long and 4 mm wide, on its lingual surface. The wear facet follows the long axis of the crown in orientation, contains homogeneously oriented striations offset from the major axis of the facet by ~15°, and has worn away edges, matching the facets on tyrannosaurid teeth described by Schubert and Ungar (2005). Though barely visible due to the presence of vertical cracks in the specimen, the tooth bears a ridge that curves toward the apex, suggesting the tooth to be from a tyrannosaur more derived than basal taxa like *Eotyrannus* (e.g., Zanno and Mackovicky, 2011; Krume-nacker et al., 2017). NJSM 16601 is nearly symmetrical labiolingually in mesial and distal views. This tooth is eroded basally on its mesial and distal ends, but nevertheless has a CBW/CBL ratio of 0.53, just under the cutoff (0.60) for the incassate condition (Brusatte et al., 2010, 2014). Nevertheless, the ziphodont nature of this specimen suggests it may belong to *Dryptosaurus* or a close relative because *Dryptosaurus* is unusual among derived tyrannosauroids in having ziphodont dentition (Brusatte et al., 2011).

Materials.—NJSM 14682, a pedal phalanx III-2 (Fig. 1.1–1.6), NJSM 16601, maxillary or dentary tooth crown (Fig. 2.1–2.5).

Remarks.—NJSM 14682 is referred to Tyrannosauroida based on the following combination of features: massive size, mediolateral width of phalanx at the distal hemicondyles equal to that at the proximal articular facet leading to a robust, proximodistally shortened morphology, a mediolaterally “pinched” diaphysis in dorsal and ventral views, a deeply concave proximal articular facet, and a deepened extensor fossa on the dorsal surface just proximal to distal hemicondyles (see below). NJSM 16601 is referred to Tyrannosauroida based on the following combination of features: size, rectangular-ovoid cross-section in labial-lingual view, gentle curvature towards the apex, and labial-lingual width that is tapered distally.

cf. *Dryptosaurus* sp. (morphotype A) Cope, 1866
Figure 2.6–2.14

Description.—The partial tooth included in NJSM 12436 is the incomplete tooth of a large theropod dinosaur, measuring 27.5 mm high. This specimen is interpreted as a shed maxillary or dentary tooth based on its size and because it lacks a root. This tooth is distinct from NJSM 13734 in having a denticle count of 12 d/5 mm basally (see below), which matches the condition in *Dryptosaurus aquilunguis* (Schwimmer, 2016). These denticles are peg-like and rectangular, strongly oriented apically. The smaller tooth included in NJSM 12436 bears interdenticular sulci on its distal carina (e.g., Currie et al., 1990; Smith, 2007; Hendrickx et al., 2015) that extend slightly onto the tooth surface and are, like the denticles, oriented parallel to the base of the tooth. These denticles are approximately twice as long mesiodistally than apicobasally. The mesial surface of NJSM 12436 is slightly recurved apically, matching the condition of NJSM 16601 and the teeth of *Dryptosaurus aquilunguis* more than that of NJSM 13734. The preserved portion of the basal end of NJSM 13734 is elliptical and suggests a labiolingually compressed, teardrop-shaped morphology consistent with the ziphodont condition. However, because of the fragmentary nature of the specimen, ziphodonty on NJSM 13734 cannot be confirmed. The mesial portion of the tooth is absent. However, NJSM 13734 is relatively well preserved compared to other specimens, suggesting the tooth was shed close to the site of its deposition (e.g., Denton et al., 2011).

NJSM 13095 is another large theropod maxillary or dentary tooth, measuring 27 mm in CH. This tooth, which is ziphodont in morphology (CBW/CBL=0.53), has been abraded to a greater extent than NJSM 13734, but still preserves distal denticles that have a density of 9.5 denticles/5mm. This denticle count corresponds closely to the denticle density reported for the apical distal carina of the maxillary teeth of *Dryptosaurus* by Schwimmer (2016). NJSM 13095 is straightened distally, with a gently curved mesial outline.

Materials.—NJSM 12436, partial maxillary or dentary tooth (Fig. 2.6–2.8), NJSM 13095, partial maxillary or dentary tooth (Fig. 2.9–2.14).

Remarks.—NJSM 12346 and NJSM 13095 are assigned to Tyrannosauroidae based on the presence of the following combination of features: size, distally tapered labio-lingual depth, gentle curvature, and widely spaced, peg-like denticles. NJSM 12346 and NJSM 13095 are assigned to cf. *Dryptosaurus* due to having a similar basal distal carina denticle count of 12 denticles/5 mm and apical distal carina denticle count of 9.5 denticles/5 mm, respectively (Schwimmer, 2016). NJSM 13095 is also ziphodont in morphology, further allying the specimen with *Dryptosaurus*. Although NJSM 12346 may be ziphodont in morphology, which would further unite the tooth with *Dryptosaurus* (e.g., Brusatte et al., 2011), the presence of this morphology on NJSM 12346 is uncertain.

Tyrannosauroidae indet. (morphotype B)
Figure 3

Description.—NJSM 13734 is the partial maxillary or dentary tooth of a large theropod dinosaur. This tooth is slightly recurved and labiolingually narrow basally. The partial tooth preserves the mesial carina and approximately one quarter of its the denticles, which are small (16 denticles/5 mm), widely spaced, peg-like, and strongly oriented mesially. The distal denticles of NJSM 13734 are much smaller than those of NJSM 12436, with 17 denticles/5mm basally. This denticle density is similar to but exceeds that of the teeth of *Appalachiosaurus* (e.g., Carr et al., 2005; Schwimmer et al., 2015; Schwimmer, 2016). The denticles in NJSM 13734 are ~2.5 times longer mesiodistally than apicobasally. Interdenticular sulci that extend slightly onto the tooth surface are also present on NJSM 13734 and are about the same dimensions as the denticles themselves. The tooth, though only partially preserved, is labiolingually compressed basally (Fig. 2.11). The estimated mesiodistal length of the base of the tooth when complete is 12–15 mm, and the preserved labiolingual width of the base is 6 mm. Thus, the tooth may have been ziphodont when complete, allying the large tyrannosauroid taxon to which it belonged with *Dryptosaurus aquilunguis* (e.g., Brusatte et al., 2011).

Material.—NJSM 13734, maxillary or dentary tooth (Fig. 3.1–3.4).

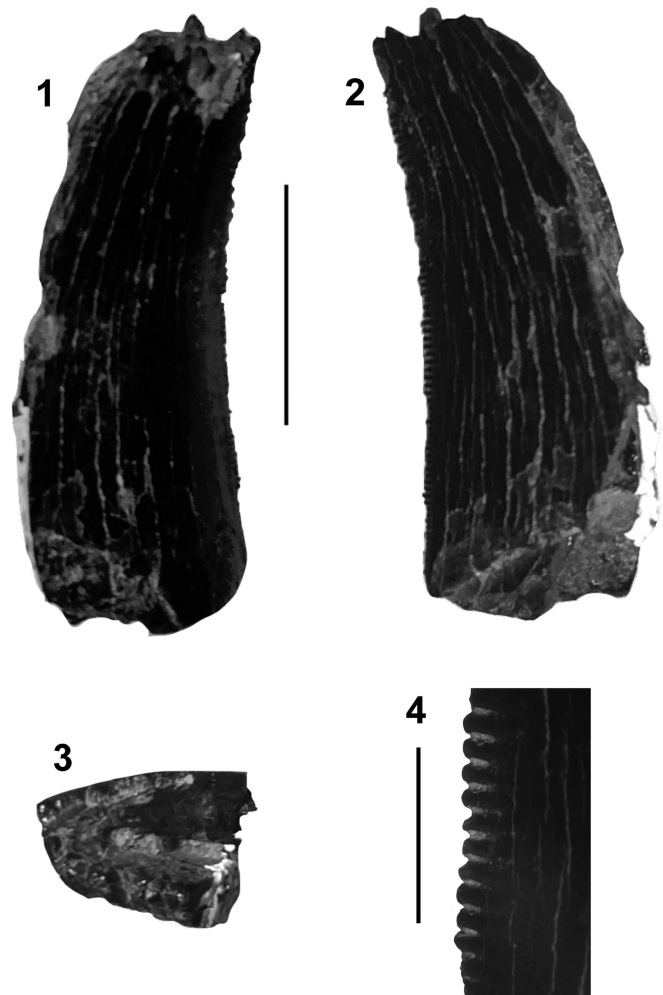


Figure 3. Partial tyrannosauroid maxillary or dentary tooth NJSM 13734 in labial (1), lingual (2), and basal (3) views, with magnified view of denticles (4). Scale bar = 10 mm (1–3), = 2 mm (4).

Remarks.—NJSN 13734 is referred to Tyrannosauroida on the basis of the following combination of features: size, having a D-shaped cross-section basally, being gently curved towards its apex and distally tapered in its labiolingual width, and possessing peg-like denticles that increase in density towards its base. NJSN 13734 is referred to a second morphotype of tyrannosauroid based on having denser denticles than the other Ellisdale tyrannosauroid teeth for which distal denticles are preserved (NJSN 12436).

Ornithomimosauria Barsbold, 1976

Ornithomimosauria indet.

Figure 4

Description.—NJSN 14686 is smaller than NJSN 14682 in all dimensions (Table 1), measuring 67 mm long proximodistally. The bone is slightly eroded, though most of the bone surface is well preserved and not noticeably abraded. Distally, the hemicondyles are separated by a very slight sulcus, which transitions to a shallow, proximally extending hyperextensor pit dorsally. The collateral ligament pits are subovoid and deep. Ventrally, the two distal hemicondyles form a concavity just proximal to the distal articular surface. In lateral and medial views, the dorsal surface is flattened and the ventral surface is gently arched. In dorsal and ventral views, the lateral and medial surfaces are also gently arched and gently curved outward towards the proximal end in expansion. The proximal end is appreciably wider than the distal (Table 1), and the proximal end bears pointed processes on its ventral corners that form a shallow concavity ventrally. The proximal end, though not noticeably eroded, lacks a vertically oriented projection, and there is no rim surrounding the proximal articular facet.

NJSN 14686 is assignable to Ornithomimosauria based on the proximally projecting ventral processes on the bone, a morphology found in both basal and derived ornithomimosaurs (e.g., Osborn, 1921, fig. 3a; Osmólska et al., 1972, fig. 17, pl. 49;

Kobayashi and Barsbold, 2005a, fig. 6.10G, 6.10H, 2005b, fig. 17B; Choiniere et al., 2012, fig. 14; Cullen et al., 2013, figs. 2, 3; McFeeters et al., 2016, fig. 11; Tsogtbaatar et al., 2017, fig. 4). The specimen's gently arched medial and lateral surfaces, flattened dorsal surface, smaller size when compared to the known pedal phalanges of eastern Tyrannosauroids, and shallow hyperextensor pit are also consistent with this assignment (e.g., Lambe, 1917, fig. 49; Osborn, 1921, fig. 3a; Osmólska et al., 1972, fig. 17, pl. 49; Carr and Williamson, 2000, figs. 3, 4, 8, 15; Holtz, 2004; Carr et al., 2005, supplemental information; Kobayashi and Barsbold, 2005a, fig. 6.10G, 6.10H, 2005b, fig. 17B; Choiniere et al., 2012, fig. 14; Cullen et al., 2013, figs. 2, 3; Farlow et al., 2013; McFeeters et al., 2016, fig. 11; Tsogtbaatar et al., 2017, fig. 4). In form, the proximally projecting processes on NJSN 14686 are not as well developed as in ornithomimids (e.g., Osborn, 1921, fig. 3a; Osmólska et al., 1972, fig. 17, pl. 49; Kobayashi and Lü, 2003, fig. 24; Cullen et al., 2013, figs. 2, 3; McFeeters et al., 2016, fig. 11; Tsogtbaatar et al., 2017, fig. 4). Rather, the specimen is more similar to the pedal phalanges III of *Harpymimus*, *Nedcolbertia*, an ornithomimosaur from the Aptian–Albian of China, and of ornithomimosaurs from the Arundel Clay of Maryland in the extent of its proximally oriented ventral processes and the ratio between its dorsoventral height and mediolateral width proximally (e.g., Table 1; Gilmore, 1920; Kirkland et al., 1998; Shapiro et al., 2003; Kobayashi and Barsbold, 2005a). Thus, it may be that the specimen represents an ornithomimosaur species basal to those found in the west during the Campanian, a condition that has been found with other dinosaur groups on Appalachia (e.g., Schwimmer, 1997; Carr et al., 2005; Brusatte et al., 2010; Prieto-Márquez et al., 2016a, 2016b). Further research into Appalachian ornithomimosaurs from the time of the Western Interior Seaway will be needed to test this hypothesis.

Material.—NJSN 14686, a pedal phalanx II-2 (Fig. 4.1–4.6).

Remarks.—NJSN 14686 is referred to Ornithomimosauria based on the following combination of features: proximally projecting ventral spurs (Fig. 4.1, 4.2), shallow hyper-extensor pit, and long form with gently proximodistally arched lateral and medial surfaces.

Maniraptora Gauthier, 1986

Dromaeosauridae Matthew and Brown, 1922

Dromaeosauridae indet. morphotype A.

Figure 5



Figure 4. Ornithomimosaur pedal phalanx NJSN 14686 in lateral (1), medial (2), dorsal (3), ventral (4), proximal (5), and distal (6) views. Arrows point to ventral spurs. Scale bar = 50 mm.

Description.—NJSN 14158 is the partial large tooth of a theropod dinosaur, measuring 25 mm in CH. The mesial and distal carinae are both preserved, though middle portion of the crown is only partially present. The tooth is clearly dromaeosaurid based on the apically recurved distal denticles on the distal carina and the extremely small size of the mesial denticles, which are rectangular. The distal denticles in NJSN 14158 differ from those in other dromaeosaurid specimens from Ellisdale in being twice as long mesiodistally than apicobasally, rectangular, and slightly ($\geq 50^\circ$) rather than heavily ($\leq 40^\circ$) inclined towards the crown apex. Small interdenticular sulci separate each distal denticle and do not extend onto the crown. The mesial denticles are very small and mesiodistally flattened,

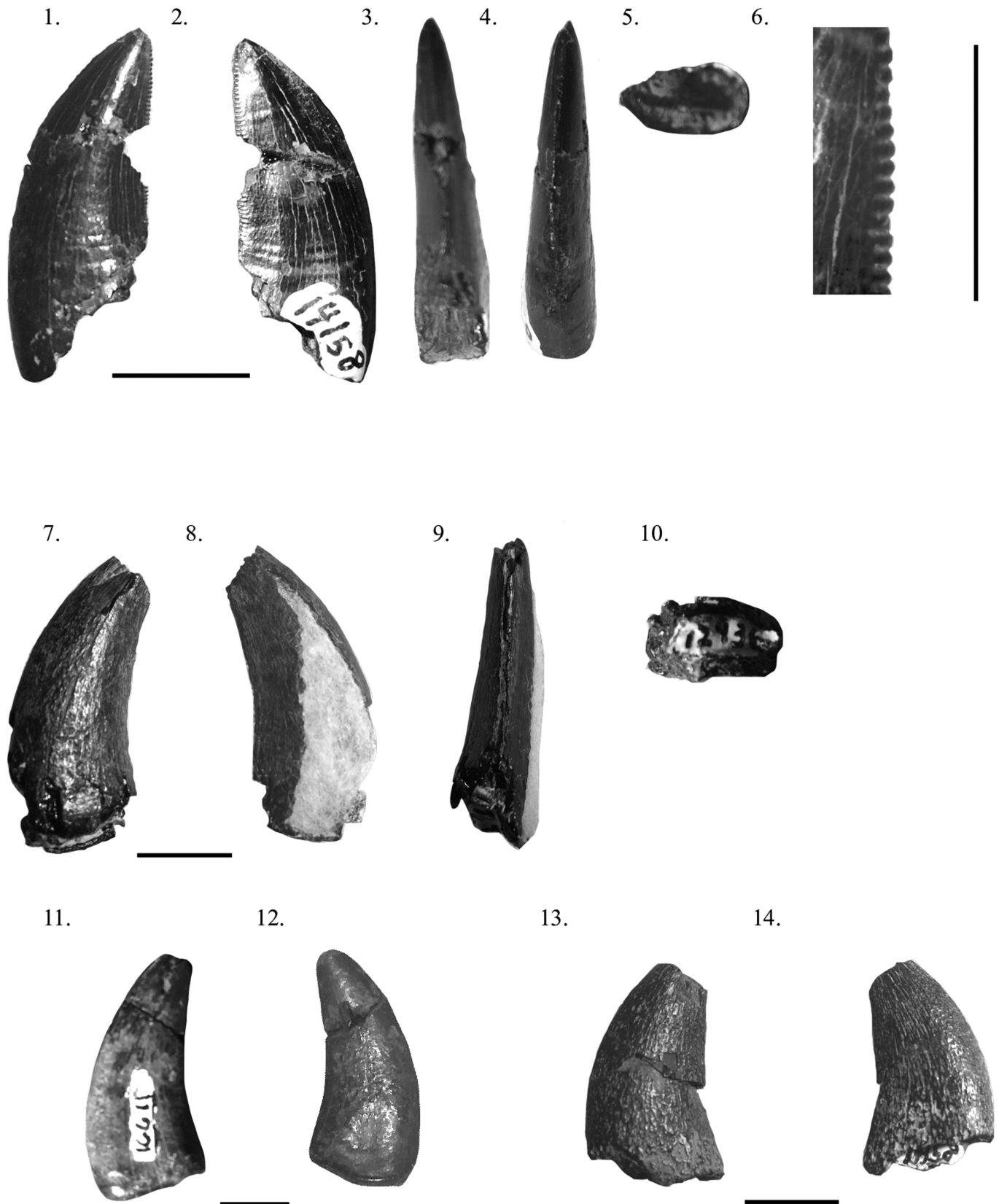


Figure 5. Dromaeosaurid maxillary or dentary teeth NJSM 14158 (larger tooth) (1–6), NJSM 12436 (complete crown) (7–10), NJSM 16611 (11, 12), and NJSM 14158 (smaller tooth) (13, 14) assigned to morphotype A in labial (1, 7, 11, 13), lingual (2, 8, 12, 14), mesial (3, 9), distal (4), and basal (5) views with magnified view of NJSM 14158 (larger tooth) (6). Scale bar = 10 mm (1–5, 7–10, 13, 14), = 5 mm (6, 11, 12).

separated by shallow interdenticular sulci. Importantly, this tooth is very large for a dromaeosaurid, falling just above the upper end of the crown height range for the giant dromaeosaurid taxon *Dakotaraptor steini* DePalma et al., 2015 and far exceeding the crown height ranges for other Late Cretaceous dromaeosaurid taxa (e.g., DePalma et al., 2015). The CHs of the Ellisdale tooth crowns are most similar to those of the teeth of *Utahraptor*, the largest known dromaeosaurid (Kirkland et al., 1993), suggesting NJSM 14158, NJSM 12436, and NJSM 16611 represent a truly gigantic dromaeosaurid taxon. The outlines of the mesial and distal faces of the tooth are recurved apically, and the specimen is laterally compressed and nearly symmetrical in distal and mesial views. In basal view, the tooth is rectangular.

The other tooth included in NJSM 14158 is heavily worn, preserving no denticles. However, the tooth is large (CH = 25 mm), ziphodont, and recurved, and is thus referred to the dromaeosaurid morphotype A. The complete tooth included in NJSM 12436 is gracile, heavily recurved, ziphodont, and large (CH = 26 + mm), representing the largest of the dromaeosaurid tooth crowns from the Ellisdale site. This latter tooth also does not preserve denticles and is D-shaped in cross-section basally.

NJSM 16611 is a dromaeosaurid tooth based on its labiolingual compression and curvature apically. Though it has been smoothed by abrasion during water transport, the size of the specimen is most similar to the large dromaeosaurid teeth included in morphotype A of the dromaeosaurid specimens of Ellisdale.

Materials.—NJSM 14158, partial tooth crowns (Fig. 5.1–5.6, 5.13, 5.14); NJSM 16611, complete tooth crown (Fig. 5.11, 5.12).

Remarks.—The tooth with preserved denticles included in NJSM 14158 is referred to Dromaeosauridae based on the following combination of features: being appreciably recurved towards the crown apex, possessing the ziphodont condition, having much larger distal than mesial denticles, and having distal denticles that curve towards the apex of the crown. The other tooth included in NJSM 14158, the eroded complete tooth crown included in NJSM 12436, and NJSM 16611 share all these features except those relevant to the denticles because none of these teeth preserves them. The morphotype represented by the tooth included in NJSM 14158 preserving denticles is characterized by large size, and peg-like denticles that are approximately twice as long as are wide and are deflected apically. The other tooth included in NJSM 14158, the complete crown included in NJSM 12436, and NJSM 16611 are tentatively referred to this morphotype on the basis of size.

Dromaeosauridae indet. morphotype B.

Figure 6

Description.—NJSM 14404 includes two tooth fragments. The larger fragment is darker in color, bearing numerous holes, which are inferred to be invertebrate borings, and enamel spalling damage on its apical end. The smaller fragment is less eroded and preserves six denticles. Each of the denticles of the smaller fragment is angled at ~45° angle with the crown towards the apex of the tooth, differentiating this specimen from the dromaeosaurid tooth morphotype that includes NJSM 14158. These denticles are approximately as mesiodistally long as apicodistally wide and are shaped like blunt hooks. The denticle density of this tooth is



Figure 6. Dromaeosaurid maxillary or dentary teeth NJSM 14404 (1) assigned to morphotype B in ?labial/lateral views. Scale bar = 5 mm.

3/mm. Both tooth fragments in NJSM 14404 are labiolingually compressed, although apicobasal curvature and the ziphodont condition cannot be confirmed in either.

Material.—NJSM 14404, tooth fragments (Figure 6.1).

Remarks.—NJSM 14404 is referred to Dromaeosauridae based on the following features: recurved, ziphodont tooth with apically oriented distal denticles.

Theropoda indet.
Figure 7

Description.—NJSM 15319 is the very small tooth of a theropod dinosaur, measuring 5 mm apicobasally. The specimen is very

recurved, more so than the tyrannosauroid but less so than the dromaeosaurid teeth from Ellisdale. NJSM 15319 preserves small, square, tightly packed denticles on its distal surface, unlike the apically recurved denticles of the dromaeosaurid or the peg-like, widely separated denticles of the tyrannosauroid teeth from Ellisdale. Mesially, the tooth is damaged and no denticles are preserved. The tooth is ziphodont, measuring 0.5 mm (CBL) and 0.24 mm (CBW) along the base. NJSM 16623 is much larger than NJSM 15319, measuring 16 mm apicobasally and representing a medium-sized theropod dinosaur. In this dimension, the tooth is slightly smaller than the smallest tooth included in the large dromaeosaurid morphotype and significantly smaller than the tyrannosauroid teeth of the Ellisdale site. The tooth is recurved apically, preserving no denticles. In basal view, the tooth is elliptical in form and ziphodont (CBW/CBL = 0.286).

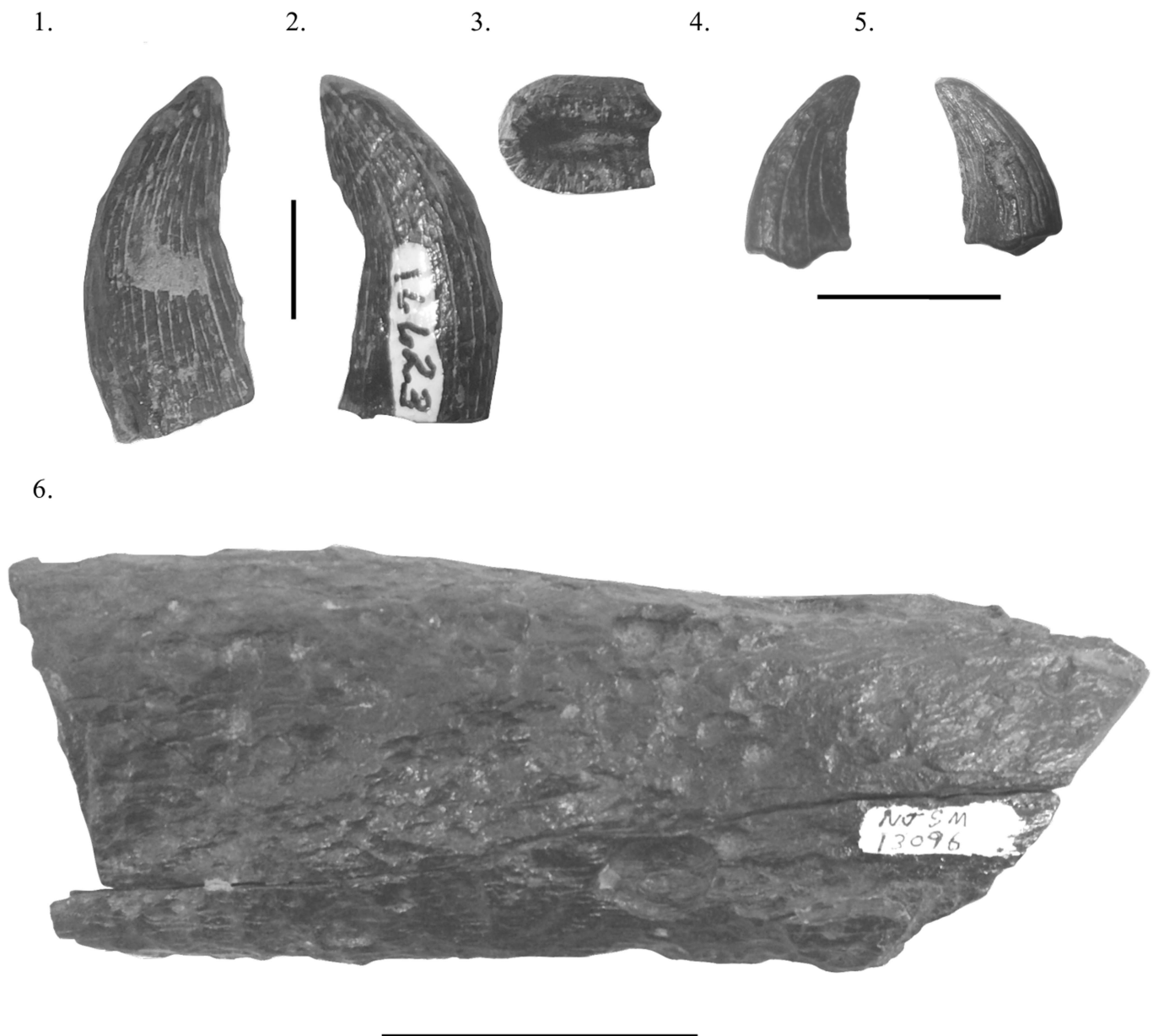


Figure 7. Indeterminate theropod specimens NJSM 16623 (1–3), NJSM 15319 (4, 5), and NJSM 13096 (6) in labial/lateral (1, 4, 6), lingual (2, 5), and basal (3) views. Scale bar = 5 mm (1–5), = 50 mm (6).

NJSM 13087, NJSM 13096, and NJSM 16607 are all partial limb shafts of theropod dinosaurs, identified as such due to their hollow interiors. NJSM 13087 and NJSM 16607 are very poorly preserved and not identifiable past large Theropoda indet. NJSM 13096, the likely partial metatarsal of a tyrannosauroid similar to *Dryptosaurus*, bears dozens of puncture marks, scrapes, and deformations from the teeth of the large crocodylian *Deinosuchus* (e.g., Schwimmer, 2002, 2010).

Materials.—NJSM 16623 (Fig. 7.1–7.3), NJSM 15319 (Figure 7.4, 7.5), maxillary or dentary teeth. NJSM 13096 (Fig. 7.6), NJSM 16607, NJSM 13087, partial limb shafts.

Remarks.—NJSM 15319 and NJSM 16623 are referred to Theropoda based on their curvature and laterally compressed state. NJSM 13096, NJSM 16607, and NJSM 13087 are referred to Theropoda based on their hollow interiors.

Discussion

Taphonomic implications of the Ellisdale theropod specimens.—The theropod specimens described vary heavily in the state of their preservation. Several specimens, including the larger tooth in NJSM 14158, the teeth NJSM 14404, NJSM 13734, and NJSM 12346, and the limb elements NJSM 16651 and NJSM 14686, are fairly well preserved, with intact denticles and other small morphological features and lacking wear from erosion. The other specimens described show much more significant wearing, such as that observed on the pedal phalanx NJSM 14682 or the smaller tooth included in NJSM 14158. These two states of preservation among the Ellisdale specimens suggest that portions of the assemblage traveled from sites varying in proximity to the final area of deposition of the bones and teeth, a model consistent with that proposed by Denton and Tashjian (2012). Indeed, the possibility remains that the Ellisdale theropod specimens represent taxa from multiple different biomes that existed progressively inland from the coastline.

Comparison of the Ellisdale and western North American theropod tooth morphotypes.—An extensive catalogue of literature has documented theropod tooth faunas from western North America, often in order to better account for dinosaur diversity and extinction during the Late Cretaceous (e.g., Currie et al., 1990; Baszio, 1997; Fiorillo and Gangloff, 2000; Sankey, 2001; Sankey et al., 2002, 2005; Weishampel et al., 2004; Fanti and Miyashita, 2009; Larson and Currie, 2013; Williamson and Brusatte, 2014). Recently, Larson and Currie (2013) quantitatively established distinct morphotypes of small theropod teeth among the specimens known from the western interior. These morphotypes include saurornitholestine, dromaeosaurine, *Paronychodon*, *Zapsalis*, cf. *Pectinodon*, cf. *Troodon*, *Richardoestes* *gilmorei*, and *R. isosceles* teeth (Larson and Currie, 2013). The latter five morphotypes are importantly distinguishable from all smallish theropod teeth collected from the Ellisdale site. None of the Ellisdale teeth have the isosceles-triangle shape characteristic of *Richardoestes isosceles*, the very large, apically oriented denticles characteristic of troodontid teeth, the prominent longitudinal ridges or blade-like mesial denticles characteristic of *Zapsalis* teeth, or the prominent enamel wrinkles

found on teeth assigned to *Paronychodon* (e.g., Larson and Currie, 2013; Williamson and Brusatte, 2014). Thus, none of the Ellisdale specimens can be referred to troodontids or *R. isosceles*. However, some of the Ellisdale teeth are somewhat similar to the saurornitholestine, dromaeosaurine, and *Richardoestes* morphotypes described by Larson and Currie (2013). As such, more explicit identification of differences between those morphotypes and the Ellisdale specimens was warranted.

The denticles of the Ellisdale theropod teeth assigned to dromaeosaurid morphotype A differ from those of saurornitholestine teeth from western North America (e.g., Larson and Currie, 2013; Williamson and Brusatte, 2014) in lacking interdenticular sulci that project onto the crown. Furthermore, the dromaeosaurid morphotype A teeth of Ellisdale are significantly larger than any of the saurornitholestine, dromaeosaurine, or *Richardoestes gilmorei* Currie, Rigby, and Sloan, 1990 teeth described by Larson and Currie (2013), their CHs are comparable to or larger than those of the teeth of the very large Maastrichtian dromaeosaurid *Dakotaraptor steini* (e.g., DePalma et al., 2015). However, the saurornitholestine teeth of western North America are similar to the Ellisdale dromaeosaurid morphotype A teeth in having apically projecting, peg-like distal denticles, a trait distinguishing both morphotypes from dromaeosaurine teeth (e.g., Sankey et al., 2002; Larson and Currie, 2013; Williamson and Brusatte, 2014). Furthermore, dromaeosaurid teeth from Ellisdale assigned to morphotype A lack the twisting mesial carina indicative of the teeth of *Dromaeosaurus* or closely related taxa (Currie et al., 1990; Currie, 1995; Sankey et al., 2002; Turner et al., 2012; Larson and Currie, 2013; Williamson and Brusatte, 2014). The Ellisdale dromaeosaurid morphotype A teeth are further distinguished from *R. gilmorei* in having relatively large, visible denticles compared to the small, minute ones of that taxon (e.g., Larson and Currie, 2013; Williamson and Brusatte, 2014).

The dromaeosaurid morphotype B teeth of the Ellisdale site are, like the teeth included in morphotype A, distinguishable from dromaeosaurine teeth in having stronger apically projecting distal denticles (forming an ~60° angle with the carina) and from saurornitholestines in not having interdenticular sulci that extend onto the tooth crown to create a “peg-like” outline for each denticle (e.g., Larson and Currie, 2013; Williamson and Brusatte, 2014). Rather, the dromaeosaurid morphotype B teeth from Ellisdale possess wave-like, “hooked” distal denticles that are most similar to those of *Paronychodon*, although teeth assigned to the latter taxon are wider labiolingually at their bases and have prominent enamel wrinkles (e.g., Larson and Currie, 2013; Williamson and Brusatte, 2014). Indeed, the distal denticles of the dromaeosaurid teeth from Ellisdale included in morphotype B are arguably more apically recurved than those of western saurornitholestine teeth and certainly more so than those of *R. gilmorei* (e.g., Larson and Currie, 2013).

As noted previously in the literature, differences in tooth morphology among tyrannosauroids are harder to quantify due to the morphological similarity of the teeth of different tyrannosauroid taxa (e.g., Samman et al., 2005; Williamson and Brusatte, 2014). However, several key differences distinguish the Ellisdale tyrannosauroid teeth from those of western tyrannosaurids. The ziphodont condition in some western tyrannosauroid teeth has been discussed as an indicator of the juvenile nature of the individuals from which such teeth came

rather than an indicator of the position of such tyrannosauroids outside Tyrannosauridae (e.g., Carr, 1999; Currie, 2003; Sampson et al., 2005; Williamson and Brusatte, 2014). However, the CHs of the Ellisdale tyrannosauroid teeth are comparable to or exceed those of subadult or adult western tyrannosaurid teeth (e.g., Carr et al., 2005, supplemental information). Thus, the presence of the ziphodont condition in likely all the tyrannosauroid teeth from Ellisdale is considered indicative that all these teeth are from large non-tyrannosaurid tyrannosauroids (e.g., Holtz, 2004; Carr et al., 2005; Brusatte et al., 2010, 2011, 2014; Loewen et al., 2013; Williamson and Brusatte, 2014; Brusatte and Carr, 2016; Carr et al., 2017). This referral is supported by the assignability of several of the Ellisdale teeth to *Dryptosaurus* or a closely related taxon and the similarity of the tyrannosauroid morphotype B teeth described to *A. montgomeriensi*s (e.g., Carr et al., 2005; Brusatte et al., 2011; Schwimmer et al., 2015).

Appalachia was at least partially isolated from Laramidia since the Albian–Cenomanian (e.g., Roberts and Kirschbaum, 1995; Russell, 1995; Schwimmer, 2002). However, it is notable that the theropod assemblage from Ellisdale does not resemble that of the Cenomanian Wayan Formation of the northern Rockies, the latter including incrassate-toothed small- and medium-sized tyrannosauroids, small dromaeosaurids, *Richardoestes*ia *isosceles*-like theropods represented by teeth, neovenatorids, and oviraptorosaurs (Krumenacker et al., 2017). The Ellisdale theropod assemblage also differs from that of the Albian–Cenomanian Mussentuchit Member of the Cedar Mountain Formation, which includes a large neovenatorid, small tyrannosauroids, theropods with teeth similar to *Richardoestes*ia and *Paronychodon*, troodontids, and dromaeosaurids (e.g., Kirkland et al., 1999; Zanno and Makovicky, 2011, 2013). The theropod assemblage of the Albian Cloverly Formation is somewhat more similar to that of Ellisdale, containing tyrannosauroids of similar phylogenetic grade to *Xiongguanlong*, the dromaeosaurid *Deinonychus*, and ornithomimosaurs (e.g., Ostrom, 1969, 1970; Zanno and Makovicky, 2011). Unlike the Cloverly, the Ellisdale theropod fauna lacks oviraptorosaurs (Makovicky and Sues, 1998) and large carcharodontosaurid dinosaurs (e.g., Kirkland et al., 1999; Weishampel et al., 2004; D’Emic et al., 2012). Because carcharodontosaurids seem to have gone extinct in the northern hemisphere sometime during the middle of the Late Cretaceous (e.g., D’Emic et al., 2012; Zanno and Makovicky, 2013) and none are known in the Campanian of the Atlantic or Gulf coastal plains (e.g., Gallagher, 1993, 1997; Ebersole and King, 2011; Schwimmer et al., 2015), the absence of that group in the Ellisdale fauna is unsurprising. Regardless, the Ellisdale assemblage and other Appalachian theropod faunas from the Campanian lack troodontids or oviraptorosaurs (e.g., Gallagher, 1993; Schwimmer et al., 2015; Schwimmer, 2016), which are fairly widespread among Campanian Laramidian formations (e.g., Weishampel et al., 2004; Gates et al., 2010, supplemental information; Larson and Currie, 2013; Longrich et al., 2013; Sampson et al., 2013a; Williamson and Brusatte, 2014; Funston and Currie, 2016; van der Reest and Currie, 2017). This absence may be due to preservation bias, because fossils of both of these groups in Laramidia are comparatively uncommon (e.g., Weishampel et al., 2004; Gates et al., 2010, supplemental information; Larson and Currie, 2013; Longrich et al., 2013;

Sampson et al., 2013b; Williamson and Brusatte, 2014; Funston and Currie, 2016; van der Reest and Currie, 2017). Nevertheless, differences between taxa of groups present in western North America and at Ellisdale indicate that Appalachia harbored a distinct theropod fauna from Laramidia, a biogeographic phenomenon previously recognized in the literature (especially for tyrannosauroids; e.g., Carr et al., 2005; Weishampel, 2006; Brusatte et al., 2011; Schwimmer et al., 2015; Schwimmer, 2016).

Comparison of the Ellisdale and southeastern North American theropod tooth morphotypes.—Faunal provincialism, though discussed extensively for Laramidian assemblages (e.g., Lehman, 1997, 2001; Gates et al., 2010, 2012; Sampson et al., 2010; Vavrek and Larsson, 2010; Loewen et al., 2013; Sampson et al., 2013a, 2013b; Thomson et al., 2013; Lucas et al., 2016), has only recently been investigated in the eastern North American fossil record of the Cretaceous (e.g., Schwimmer, 2016). As noted in the case of Laramidian assemblages, microfossil sites have the potential to inform studies of such biogeographic patterns among vertebrates (e.g., Sankey, 2008; Gates et al., 2010; Williamson and Brusatte, 2014), and the Ellisdale site is certainly of importance in testing for the presence of this phenomenon on Appalachia in revealing a northern Appalachian theropod fauna.

In the southeastern United States, the remains of at least three theropod dinosaur genera representing three different clades have been recovered from deposits of similar age as the Marshalltown Formation (75.7–71.2 Ma; Miller et al., 2004) at Ellisdale (Campanian, palynological data indicates an age range of 76.4–79.6 Ma; Denton and Tashjian, 2012): *Appalachiosaurus montgomeriensis* from the Demopolis Chalk (Middle Campanian; e.g., Ebersole and King, 2011) and, along with cf. *Saurornitholestes langstoni*, an additional indeterminate dromaeosaurid, and ornithomimosaurs, from the Coachman Formation of northern South Carolina and *Dryptosaurus*, ornithomimosaurs, and other small theropods from the temporally equivalent and geographically nearby Tar Heel Formation of southern North Carolina (78.7–74.5 Ma; e.g., Self-Trail et al., 2004; Harris and Self-Trail, 2006; Schwimmer et al., 2015). The close proximity and equivalent age of the Tar Heel and Coachman show that the dinosaur fauna of the Carolinas during the Campanian represented a mixing of northern (e.g., *Dryptosaurus*) and southern (e.g., cf. *Saurornitholestes langstoni*, *Appalachiosaurus montgomeriensis*) forms in the formations (e.g., Schwimmer, 2016).

The Tar Heel and Coachman formations’ theropod assemblage share the presence of two tyrannosauroids, one *Dryptosaurus*-like and one *Appalachiosaurus*-like, and ornithomimosaurs with Ellisdale. However, the teeth from the Coachman Formation assigned to cf. *Saurornitholestes langstoni* are significantly smaller than the Ellisdale dromaeosaurid teeth included in morphotype A and, unlike the Ellisdale teeth included in morphotype B, have peg-like distal denticles with interdenticular sulci that project onto the crown surface (Schwimmer et al., 2015). Furthermore, teeth of possible dromaeosaurine origin (based on their curvature and large mesial denticles) reported from the Coachman Formation by Schwimmer et al. (2015) further distinguish that theropod fauna from that of Ellisdale. Overall, the differences between the theropod faunas of southeastern North

America and Ellisdale, especially among the maniraptoran taxa, support the hypothesis that faunal provincialism existed on Appalachia (e.g., Schwimmer, 2016).

Conclusions

The Ellisdale theropod fauna includes tyrannosauroids represented by at least two tooth morphotypes, small and large dromaeosaurids, ornithomimosaurs, and indeterminate theropods, and is distinct from Campanian western North American theropod faunas, much more so resembling mid-Cretaceous theropod faunas from western North America and thus supporting the hypothesis that Appalachia was a refugium for relict dinosaur clades. Furthermore, the distinction of the Ellisdale assemblage from southeastern North American theropod faunas may support the presence of dinosaur provinces on Appalachia, a biogeographic phenomenon seen in Laramidian vertebrates.

Acknowledgments

I thank D. Parris for allowing me to access the collections of the NJSM and R. Denton for providing information on the Ellisdale site and for our discussions on Appalachian faunas generally.

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Accepted 9 May 2018